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Meanwhile, low-level moist air off the Atlantic was streaming into the developing storm over the Tennessee Valley from the east-northeast. The moist air was lifted along the arctic front, freezing into all those countless trillions of dendrites.

15:1

Mid-to-late Tuesday morning, once temperatures at the surface fell below freezing, the character of the snow changed. A brief period of moderate-sized, semi-wet flakes quickly morphed to granular powder. When meteorologists attempt to predict snow accumulation (to the very inch), we must decide on a conversion factor between snow and its liquid water equivalent.

The reason for this is the models output total precipitation as inches of liquid. The commonly used 10:1 snow:liquid ratio (10 inches of snow for every one inch of liquid precipitation) is predicated on a layer of freshly fallen snow, which by volume contains much more air than actual water.

However, a layer of wet flakes that clump into a sticky snow has greatly reduced air fraction, dropping ratios to as low as 5:1. To get such a wet snow, you need either temperatures very close to melting (o C) and/or exceptionally high cloud liquid water content (i.e. buku water vapor rising into the cloud layer).

In air far below freezing, and in a moisture-limited storm, the manufactured snow is going to pile up quite dry and fluffy. CWG meteorologists realized that this would be the case for Tuesday's storm, and 15:1 became the ratio du jour for our snow accumulation predictions. A 15:1 ratio was also adopted by both the National Weather Service and Weather Prediction Center for its forecasts.

Choosing a 20:1 ratio was probably too dry for this storm. This high ratio is more typical of extremely moisture-starved storm systems, such as Alberta Clippers crossing the Upper Midwest. 20:1 produces an extremely dusty, talcum-like snow.

Currents of rising air

In spite of subfreezing cold and moisture, without rising air, there would be no clouds to manufacture snow. In Tuesday's storm, ascending air came from three sources: The jet stream, the developing coastal low, and forced uplift along the nose of the arctic front.



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Figure 3: Dynamical processes in the polar jet stream leading to formation of coastal low (red ‰) and its intensification. Note the change in orientation of the shortwave trough (white dashed line). Adapted from Unisys Corp.

On account of some fairly advanced fluid dynamics, air ascends just ahead of the axis of a shortwave trough. Like a vacuum, the trough draws up air from the surface, dropping the pressure in the center of the low.

But the character of the shortwave changed through the day on Tuesday. The tilt of the trough axis became re-oriented from northeast-southwest to northwest-southeast. When a trough acquires this "negative tilt" configuration, air is exhausted from the top of the surface low much more rapidly. As a result, the low began to intensify along the coast Tuesday afternoon through the overnight. The heaviest snow tends to fall during these intensification phases. This was consistent with the burst of heavy snow across our region around 6 p.m.

Snowband

All the elements that conspire to create heavy snow – cold air and ascending moisture – tend to concentrate into band or streak-like features on the backside of a developing vortex. Predicting the occurrence, location, timing and duration of bands is the essence of a snowstorm's nowcasting phase.

Meteorologists know about the preferred regions for these bands, and you may hear jargon such as "frontogenetic region" or "deformation zone". Fluid dynamics aside, Tuesday's storm had all this. Figure 4 shows a very prominent, elongated band that developed during the late afternoon on Tuesday – stretching from Washington to Boston.



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